METEOR-Berichte

# *Submarine volcanism in the western Sicilian Channel*

Cruise No. M191 (GPF 21-1\_036)

16.07.2023 – 05.08.2023, Algeciras (Spain) – Piraeus (Greece) SUAVE



**Jörg Geldmacher, Aaron Micallef, Odysseas Archontikis, Meret Felgendreher, Giulia Matilde Ferrante, Jonathan Ford, Jacqueline Grech Licari, Folkmar Hauff, Silke Hauff, Alastair Hodgetts, Jakob Lang, Kerys Meredew, Christian Timm, Maxim Portnyagin, Sebastian Watt**

Chief Scientist: Jörg Geldmacher GEOMAR Helmholtz Centre for Ocean Research Kiel



#### **1 Cruise Summary**

#### **1.1 Summary in English**

The origin and role of volcanism in passive continental rifts remains poorly understood relative to other volcano-tectonic settings. The western Sicilian Channel in the central Mediterranean Sea represents an area of pronounced crustal extension with a variety of volcanic landforms closely associated with extensional faults. The main goal of our study is to analyze how volcanism has developed in conjunction with tectonic structures in the western Sicilian Channel. Therefore, we have extensively mapped (with multibeam echo sounder, sediment echo sounder, and towed magnetometer) the seafloor and sampled (chain bag dredge) unexplored volcanic structures along the Sicilian Channel Rift Zone and the Capo-Granitola-Sciacca Fault Zone. One unexpected result is that many features shown in the predicted bathymetry as seamount-like elevation, and which were previously interpreted as presumably volcanic edifices/cones, turned out to be either nonexistent or no evidence for an igneous nature (magnetometer data/dredging results) could be established. On the other hand, three previously unknown volcanic outcrops were discovered and partly sampled. The obtained data and collected samples will provide new insights into the role of magmatism in regions of continental extension, and will allow us to develop a tectonic and magmatic framework for the western Sicilian Channel.

#### **1.2 Zusammenfassung**

Der Ursprung und die Rolle des Vulkanismus in passiven kontinentalen Dehnungsgebieten ist im Vergleich zu anderen vulkanisch-tektonischen Umgebungen noch immer wenig verstanden. Der westliche Sizilianische Kanal im zentralen Mittelmeer ist ein Gebiet mit ausgeprägter Krustendehnung und einer Vielzahl von vulkanischen Bildungen, die eng mit Dehnungsstörungen verbunden sind. Das Hauptziel des Vorhabens besteht darin zu analysieren, wie sich der Vulkanismus in Verbindung mit tektonischen Strukturen im westlichen Sizilianischen Kanal zeitlich und kompositionell entwickelt hat. Daher haben wir den Meeresboden entlang der Straße von Sizilien-Grabenzone und der Capo-Granitola-Sciacca-Verwerfungszone ausgiebig kartiert (Mehrkanal-Echolot und geschleppten Magnetometer) und bisher unbeprobte Vulkanstrukturen beprobt (Kettensackdredge). Ein unerwartetes Ergebnis ist, dass viele Strukturen, die in der vorhergesagten (Satelliten-) Bathymetrie als kegelförmige Erhebungen angezeigt werden, und die bisher als vulkanische Zentren interpretiert wurden, sich entweder als nicht existent herausstellten oder keine Beweise für eine vulkanische Entstehung (Magnetometerdaten/Dredgeproben) erbracht werden konnten. Andererseits wurden drei bisher unbekannte vulkanische Strukturen neu entdeckt und zum Teil beprobt. Die gewonnenen Daten und Proben werden daher neue Erkenntnisse über die Rolle des Magmatismus in Regionen kontinentaler Ausdehnung liefern und es uns ermöglichen, ein tektonisches und magmatisches Modell für den westlichen Sizilianischen Kanal zu erstellen.

# **2 Participants**

## **2.1 Principal Investigators**



## **2.2 Scientific Party**



## **2.3 Participating Institutions**





The M191 scientific party

## **3 Research Program**

#### **3.1 Description of the Work Area**

#### **3.1.1 General Introduction**

Continental rifting gives rise to thinned lithosphere through extensional tectonics and is often associated with volcanism. The origin of this volcanism remains poorly understood relative to other volcano-tectonic settings on Earth. Furthermore, the role of mantle decompression, melting and magmatism within rifting processes is debated and may vary between settings. Active rifting, where mantle upwelling and associated magmatism drive lithospheric extension (e.g. Kendall et al., 2005), represents one endmember process. Volcanism also occurs in passive extensional settings (Franke, 2013; Putirka & Platt, 2012), where it is interpreted to occur through local mantle decompression melting in response to crustal thinning. Knowledge of active rifting processes has been greatly advanced in regions such as Afar, Ethiopia (Armitage et al., 2015; Ferguson et al., 2010), but there remain uncertainties about the precise nature and role of volcanism in rifts that are not thought to be associated with mantle plume ascent.

#### **3.1.2 The Sicilian Channel**

The Sicilian Channel is a shallow water platform (<400 m in depth) in the central Mediterranean Sea, located between the southern coast of Sicily, eastern coast of Tunisia and the Malta Escarpment (Figure 3.1). It forms part of the northern African continental plate, thinned by extension since the late Miocene (Civile et al., 2008), and comprises a Plio-Quaternary to lower Messinian clastic sequence that overlies 6-7 km of Mesozoic-Cenozoic shallow to deep-water carbonate deposits with intercalated volcanic rocks. The submarine landforms of the Sicilian Channel include shallow continental shelves, morphological banks of sedimentary or volcanic origin, fault-controlled rift basins, and a foredeep basin. The Sicilian Channel has been characterised by complex geodynamic processes and the co-existence of multiple stress fields in space and time. There are three key tectonic phases in the western Sicilian Channel that led to the formation of the following, broad structural elements: (i) A Neogene ~NNW-SSE front of a continental collision between the African and Eurasian plates, which led to the closure of the Tethyan ocean and gave rise to the Sicilian-Maghrebian thrust belt in the north-western (Sicilian-Maghrebian thrust front) and northern part (Gela Nappe) of the Sicilian Channel. (ii) A Pliocene-Quaternary rift-related zone in the central part of the Sicilian Channel, composed of three NW-SE oriented grabens (Pantelleria, Linosa and Malta) bounded by parallel, sub-vertical normal faults. (iii) The Capo-Granitola-Sciacca Fault Zone (CGSFZ) (Civile et al., 2018), an up to 40 km wide and 70 km deep, NNE-oriented lithospheric shear zone that extends for at least 200 km from the southern coast of Sicily to Linosa Island (Calo & Parisi, 2014; Caracausi et al., 2005; Civile et al., 2018; Civile et al., 2014; Fedorik et al., 2018; Ghisetti et al., 2009).

These geodynamic processes have given to rise to a range of volcanic manifestations at the surface. The volcanic activity in the western Sicilian Channel is predominantly anorogenic. Magmatic features include the islands of Pantelleria and Linosa, several submarine volcanic edifices (e.g. Graham, Terribile, Adventure, Nameless Banks), and buried dykes and sills (Civile et al., 2018). Although some magmatic products date to 9.5 Ma (Nameless Bank; (Beccaluva et



al., 1981)), the other four volcanoes that have been dated have a Pliocene/Pleistocene to recent age (Calanchi et al., 1989; Rotolo et al., 2006). Submarine eruptions have been reported up to historical times.

**Fig. 3.1:** Bathymetric map of the Sicilian Channel with main structural and geomorphic features labelled. PG: Pantelleria Graben; MG: Malta Graben; LG: Linosa Graben; GN: Gela Nappe; AB: Adventure Bank; GB: Graham Bank; TB: Terribile Bank; NB: Nerita Bank; NaB: Nameless Bank; PB: Pinne Bank; SMTF: Sicilian-Maghrebian Thrust Front: ME; Malta Escarpment; CGSFZ: Capo-Granitola-Sciacca Fault Zone: CGFS: Capo-Granitola-Sciacca Fault System; SFS: Sciacca Fault System.



**Fig.3.2:** Overview map of location of dredge stations within main working area (encircled by black stippled line) with graphical symbolization of recovered material and ships path (white line).

## **3.2 Aims of the Cruise**

In this project we will investigate how volcanism has developed in conjunction with tectonic structures in the western Sicilian Channel in order to propose a tectonic and magmatic framework for the region and improve our understanding of the role of magmatism in regions of continental extension. Specifically, the scientific objectives of the cruise were to map and sample the unexplored submarine volcanic manifestations along the CGSFZ and the Pantelleria and Linosa Grabens (Fig. 3.2). The collected data will be used to:

i. Determine the age, mantle source, melting process and crustal evolution of magmas across the area, and interpret these alongside data from volcanic centers in the central Mediterranean (Etna, Hyblean Plateau, Ustica, Alicudi) to evaluate mantle melting processes in a regional tectonic framework,

ii. Discriminate the volcanic signatures (melting, ascent and timing) of the magmatic manifestations at the seafloor between extensional and strike-slip kinematic fields,

iii. Assess the relationship between the architecture and kinematics of the CGSFZ and the Sicily Channel Rift Zone, and the spatial distribution of volcanic activity,

iv. Infer the structural framework and stress field in the southern section of the CGSFZ and its spatial and temporal evolution,

v. Assess the type, extent and age of volcanic activity across the region to evaluate present-day hazards.

## **3.3 Agenda of the Cruise**

## (J. Geldmacher)

To achieve the scientific goals of the SUAVE research project, cruise M191 conducted area-wide as well as more targeted multi-beam mapping, sediment echo sounding and magnetometer surveys, and carried out rock sampling at structures identified as possible igneous edifices or exposed lava fields/domes. Very few of these structures have been mapped or sampled before. Regarding specific working areas, the following structural features were selected for closer investigation:

- Bathymetric highs along the N-S trending Capo-Granitola-Sciacca Fault Zone (CGSFZ), including a volcanic cone field near the Sicilian coastline, Galatea shoal, Graham and Terribile Banks, Cimotoe Seamount, Pinne Bank, and the volcanic cone field around Linosa Island.
- NW-SE striking graben structures in the central Sicilian Channel (Pantelleria-, Linosa- and Malta-Graben).
- Presumable volcanic features to the east (Nameless Bank, Madrepora) and west (Pantelleria region) of the CGSFZ, which elongated outlines resemble the NW-SE direction of the graben structures. Unfortunately, no permission for work around Pantelleria was granted by the territorial authorities.

## **4 Narrative of the Cruise**

## (J. Geldmacher)

All 15 members of the scientific party arrived in Algeciras in the afternoon/evening of July 14 and boarded the vessel the next morning on July 15. The GEOMAR equipment container and the OGS magnetometer were already loaded and placed on the working deck. The vessel left the port of Algeciras on the following day, July 16 at 09:00, to start its almost 4 days transit to the main working area in the Sicilian Channel.

The transit time was used to unpack the containers, set up the laboratories, test the instruments, and to conduct other preparations for the cruise (including safety drills and rehearsal of the lab workflow). While transiting through the Spanish EEZ, surface water samples (for nannoplankton research) were collected approximately every 4 h. The water sampling continued at similar sequences in Italian and Greek waters during the entire time span of the expedition.

In the early morning on July 19, we arrived in the Italian EEZ and the EM 120 and EM 710 multibeam echo sounders (MBES) and sediment echo sounder (PARASOUND) were switched on. Shortly afterwards an SVP (sound velocity profile) probe was deployed in 1900 m water depth (to calibrate the MBES data). While the ship was stationary, the magnetometer was brought into the water to test all components of the system. Subsequently, the vessel accelerated and continued its transit to the Sicilian Channel. After a few hours, the magnetometer was recovered again.

On the morning of July 20, we arrived in the first working area located between the Sicilian coast and Graham Bank, along the northern rim of the Sicilian Channel (Fig. 3.1). Since the water is much shallower here (generally < 300m), another SVP was acquired before a systematic mapping survey (with multibeam echosounder, PARASOUND and magnetometer) was conducted. The survey covered all known submarine cone-shaped structures (believed to be of volcanic origin). Dredge site locations were identified and it was assured that each dredge track was placed at a sufficient distance from the nearest underwater cable This evaluation followed a standardized procedure applied for all dredge sites during this expedition (criteria: >1000 m and >2.5 times water depths distance from the nearest cable). In total, eight dredge hauls (M191-5,-6,-7,-8,-10,-11,-12,-13,-14) were conducted in this area in the course of July 20 and the following day, June 21. During the night, an extensive mapping and magnetometer survey was conducted in an area where no multibeam/magnetic data had been acquired before (M191-9). In the afternoon of July 21, the vessel transited southwards towards Graham Bank, where a short (c. 2 h) mapping survey (multibeam and magnetometer) was run over all sites (cones) of potential interest for dredging (M191-15), Two dredge hauls were conducted on one volcanic cone (M191-16,-17) near Graham Bank's northern edge. The summit of Graham Bank consists of two prominent cones with flat tops, the southernmost of which rises up to 9 m below sea level. This shoal, well known to local fishermen, represents the site of the most recent volcanic eruption in the working area. In July 1831, a volcanic outburst produced a short-lived island ("Ferdinandea Island") that was washed away by wave erosion a few months thereafter. According to historical records, such emergence/disappearance occurred already four times in this area since 300 BC. Accordingly, we had high hopes to obtain fresh volcanic rocks from this site. The night of July 21/22 was spent with another extensive mapping survey (M191-18) comprising the area SW of Graham Bank and the cluster of prominent cones between Cimotoe and Pinne Bank.

On the morning of July 22, another VSP was conducted (M191-19) before dredging on/near Graham Bank resumed (M191-20 to -24). At Site M191-20, active venting ("flares") was observed in the water column sonar data (Fig. 4.1) and the dredge haul recovered several tube-shaped rocks interpreted as vent chimneys (Fig. 4.2). After a short mapping/magnetometer survey (M191-25), the last dredge haul of the day targeted one of the cones between Cimotoe and Pinne Bank (M191- 26) before we spent the night carrying out further mapping/magnetometer surveying in the area around the Galatea shoal (M191-27).







**Fig.4.2**: Scientist Kerys Meredew from the University of Birmingham (UK) holding a chimney structure, dredged from near the active venting site of Graham Bank (see Fig. 4.1) (Photo: J. Geldmacher).

On July 23, four sites were dredged in the newly mapped area (M191-28,-29,-30-31), none of which returned volcanic material. Since we experienced a decrease in the quality of the multibeam data, another SVP was conducted before the day was concluded with an extensive mapping/magnetometer survey (M191-33) that brought the vessel back to Graham Bank. Dredge hauls M191-34,-35,-36 were conducted along a previously unknown chain of N-S- striking cones west of the Graham Bank summit. Near the northernmost cone of Graham Bank another vent field was observed in the water column sonar data. Afterwards, we spent the remaining time of July 24 dredging at and near the summit of Terribile Bank (the eastern neighbour of Graham Bank), where magnetic and backscatter data implied the presence of igneous and exposed rocks (M191-37 to - 45). However, most dredges returned carbonate crust and only one volcanic rock was recovered (site M191-37).

The night and most of the following day, July 25, was spent carrying out a more detailed mapping/magnetometer survey of Terribile Bank and the area NE and SW of Pinne Bank (M191- 46). During this survey, intense flare activity was observed in the water column at 37°08'N, 13°11' indicating large-scale gas venting. Eventually, dredging resumed along the NE flank of Nameless Bank (a few miles east of Pinne Bank). Once more, the positioning of the dredge sites was guided by identified magnetic anomalies, record during the preceding mapping survey. Dredge hauls M191-47, -48,-49 were conducted up the steep NE slope of Terribile Bank (returning only consolidated limestone) before the vessel turned southwards. The night was devoted to mapping/magnetometer surveying (M191-50).

In the morning of July 26, the vessel arrived in our southernmost working area, north of the island of Linosa and the Linosa Graben. Immediately north of Linosa, numerous steep volcanic cones /seamounts were found. However, like in all previous working areas, the choice of potential dredge

sites was greatly restricted by the presence of several underwater cables. In addition, the environmental protection zone around Linosa prevented all activities within approximately 2.5 nm from the shoreline (which included several volcanic cones). Dredging of the only two cones that were accessible to us (M191-51, -52, and -54 to -57), however, retrieved excellent lava samples (Fig. 4.3). The night and the following morning on July 27 were again used for mapping/magnetometer survey. No further volcanic cones were found along the northern slope of the Linosa Graben but a distinct magnetic anomaly indicated that that the large solitary seamount on the carbonate platform between the Linosa and Malta Grabens might contain igneous rock. Two dredge hauls (M191-59, -60) were carried out at the lower and upper slope of this structure.





The night and most of the following day, July 28, was devoted to multibeam/magnetometer mapping the area north of the Linosa graben (M191-61) but no structures that could have been identified as of volcanic origin were found and the vessel returned to the dredging area that was sampled the day before to obtain more igneous material from this edifice (dredge hauls M191-62) to -66). The subsequent transit back to the northern part of the Sicilian Channel was used for multibeam/magnetometer mapping (M191-67).

On July 29, we dredged on the summit area of Terribile Bank (dredge hauls M191-68 to -73). This shoal is well-known by local fishermen and the selection of dredge tracks was therefore partly hampered by the intense fishing activity. Despite this challenge, dredging of a small pockmark field (M191-73) was particular successful and returned interesting igneous rocks. Further multibeam/magnetometer around Terribile Bank was conducted during the night (M191-74).

On July 30, dredging proceeded temporarily to Nameless Bank (around 20 nm SE of Terribile Bank) with dredge hauls M191-75, and -76, before mapping (M191-77) and dredging resumed on Terribile Bank (M191-78, -79), where fishing activity had decreased significantly in the meantime.

A further multibeam/magnetometer survey (M191-80) brought the vessel back to neighboring Nameless Bank on July 31, where two dredge hauls were conducted near its summit (M191-81, - 82). The day was concluded by successfully dredging the upper southeastern slope of the Bank (M191-83,-84,-85). While mapping on Nameless Bank, an elongated, c. 100 m x 16 m long feature was found lying on the seafloor in about 100 m water depth, which looked like a ship wreck on the backscatter data.

The following day, Aug. 1 was devoted to further systematic multibeam/magnetometer mapping of the southern portion of Nameless Bank (M191-86) and four dredge hauls were subsequently carried out on the eastern part of this bank (M191-87 to -90). The last night of this expedition in the main working area was spent on multibeam/PARASOUND/magnetometer survey of the northern flank of the Malta Graben (M191-91), where several seamount-like structures were mapped. Since no magnetic anomaly was found, the search during the following morning, Aug. 2, focused on a vast bathymetric high located halfway between the Malta Graben und Nameless Bank. However, no further magnetic anomalies were found. After the final recovery of the magnetometer, a last SVP (M191-92) concluded our research in the main working area of the Sicilian Channel. Underway multibeam mapping, however, continued until the vessel left the Italian EEZ on Aug. 3. While most scientists spent the 2.5 days transit time cleaning the labs, packing the equipment and report writing, surface water sampling for nannoplankton research continued until we reached the port of Piraeus. At 08:35 on Aug. 5 the first line was fixed to the pier and Exp. M191 officially ended. In the afternoon all scientists disembarked the vessel.

In total, 18 dedicated multibeam/magnetometer surveys, covering a total length of 2886 km, were conducted to record high-resolution bathymetric data, backscatter, water column profile data and the total magnetic field intensity. We have carried out dredging at 68 stations, only 11 of which returned empty or contained just unconsolidated mud. All others delivered carbonate crusts, consolidated limestone (most likely from the continental carbonate platform) and, at 24 dredge stations, igneous rocks. In addition, 82 surface water samples for nannoplankton research were taken along our track between Algeciras and Piraeus. No deployed device was lost or seriously damaged.

## **5 Preliminary Results**

#### **5.1 Bathymetric Mapping and Hydroacoustics**

(A. Micallef)

#### **5.1.1 System Overview and Data Processing**

#### a) Multibeam echosounder

RV METEOR is equipped with two Kongsberg Maritime multibeam echosounder (Fig. 5.1). The EM122 system operates at 12 kHz and covers water depths from 20 meters below the transducers up to full ocean depth; while the EM710 system offers a frequency range from 70-100 kHz of signals for water depths ranging from 3 m below transducers to roughly 1000 m. Two different transmit pulses can be selected: a CW (Continuous Wave) or FM (Frequency Modulated) chirp.

The sounding mode can be either equidistant or equiangular or mixed, depending on operation preferences and requirements. Both systems can be operated in single-ping or dual-ping mode, where one beam is slightly tilted forward and the second ping slightly tilted towards the aft of the vessel. The whole beam can also be inclined towards the front of the back and the pitch of the vessel can be compensated dynamically. The EM122 system produces 432 beams covering a swath angle of up to 150° while the EM710 system produces 432 beams for a maximum swath angle of 140°. Both systems offer a high-density beam-processing mode with up to 800 soundings per swath. The swath angle, however, can be reduced, if required. The transducers of both multibeam echosounder systems of RV METEOR are mounted in a so-called Mills cross array, where the transmit array is mounted along the length of the ship and the receive array is mounted across the ship. The system on RV METEOR is of a 1° x 2° design. The EM710 system installed on RV METEOR is of a 1° x 1° design, but transducers are much smaller. The echo signals detected from the seafloor go through a transceiver unit (Kongsberg Seapath) into the data acquisition computer or operator station. In turn, the software that handles the whole data acquisition procedure is called Seafloor Information System (SIS). In order to determine the point on the seafloor where the acoustic echo is coming from, information about the ship's position, movement and heading, as well as the sound velocity profile in the water column are required. Positioning is implemented onboard RV METEOR with conventional GPS/GLONASS plus differential GPS (DGPS) by using either DGPS satellites or DGPS land stations resulting in quasi-permanent DGPS positioning of the vessel. These signals also go through the transceiver unit (Seapath) to the operator station. Ship's motion and heading are compensated within the Seapath and SIS. Beamforming also requires sound speed data at the transducer head, which is available via a sound velocity probe. This signal goes directly into the SIS operator station. Finally, a sound velocity profile for the entire water column can be obtained either from a sound velocity probe or from a CTD (conductivity, temperature and density) probe. During cruise M191, we used direct sound velocity measurements with a special profiler probe at various stages of the cruise.



**Fig. 5.1** Schematic sketch illustrating the principal mode of operation of multibeam echo-sounding systems. The whole angular coverage sector  $(\alpha)$  of the KONGSBERG EM 122 system is up to 150°.

In addition to bathymetric information, both the EM122 and the EM710 system register the amplitude of each beam reflection as well as a sidescan signal for each beam (so-called snippets). Both systems also allow recording the entire water column. The amplitude signals correspond to the intensity of the echo received at each beam. It is registered as the logarithm of the ratio between the intensity of the received signal and the intensity of the output signal, which results in negative decibel values. For each ping, both the EM122 and the EM710 record 432 backscatter intensity values. The water column data correspond to the intensity of the echoes recorded from the instant the output signal is produced. All echoes coming from the water column, the seabed and even below the seabed are recorded for each beam. When the water column data of one ping is divided into a starboard and portside subset, one can produce two traces, one for each subset. Each trace is built up as a time series in which for each time the highest amplitude is selected from all beams. Then the starboard and the port traces are joined together.

During cruise M191 the following settings of the Kongsberg EM122 system were used. The pulse was FM, ping mode was set to HD-equidistant, dual ping mode was set to dynamic, and depth mode was set to automatic. The beam angle was 140° during most of the survey. Survey speed varied between 4 and 10 knots. Data were acquired continuously, except when the ship was stationary during dredging operations and sound velocity profiling. Acquisition parameters for the EM710 system were the same as those for the EM122. Water column data were recorded during most of the surveys using the EM710 only. Six sound velocity probe casts were used for water sound velocity profiles (Figure 5.2).



Fig. 5.2: Location of conducted sound velocity profiles (SVP) within the Italian EEZ.

Data processing was carried out onboard using QPS Qimera and FMGT software. After loading the raw data (.all files) and the correct sound velocity profile, a dynamic surface was created showing the ship's track and the raw data. Qimera allows an automatic elimination of major erratic data points using a spine filter. Furthermore, there are several tools for detailed elimination of

erratic data points, for example a swatch editor, a 2D editor or a 3D editor, which enable the operator to process each single beam stepwise. All editors display not only the cleaned data but also, if desired, the rejected data points and offer a variety of visualizations of the data (according to files, depth, intensity etc.). After data cleaning, a static surface was generated from the dynamic surface creating a .sd file, which was loaded in the QPS Fledermaus software, allowing 3D visualization of the cleaned data. Furthermore, the data can be imported in FMGT to generate backscatter grids. Here, radiometric corrections, filtering, angle-varying gain and anti-aliasing filters and topographic corrections were applied to the backscatter data before outputting a georeferenced mosaic. To deliver an immediate 3D impression of the bathymetry, also uncleaned data were visualized with QPS Fledermaus and used for the quick selection of dredge tracks.

Both the EM122 and the EM710 multibeam echosounder produce a second type of raw data files with extension \*.wcd, which stores water column data. These files were imported and viewed in Qimera.

#### b) Sediment echo sounder

RV METEOR is equipped with an ATLAS PARASOUND P70 sediment echo-sounding system. Sediment echo-sounding systems (or sub-bottom profilers) are used to image sub-seafloor geological structures such as, for example, marine sediment successions. Within the survey area the system was mainly used for analysis of sedimentary processes, such as the identification of mass transport deposits, background sedimentation, and tectonic surface deformation. In addition, the sub-bottom profiles discriminate well between volcanic and sedimentary material, as volcanic material strongly attenuates the signal and leaves vertical "blanking" zones. For the standard operation we chose a primary high frequency (PHF) of 22 kHz and a primary low frequency (PLF) of 4 kHz, which implies a secondary high frequency (SHF) of 48 kHz and a secondary low frequency (SLF) of 4 kHz. The source pulse was a continuous rectangular wavelet. The receiver amplification was -5 dB gain (PHF) and 20 dB gain (SLF). Both the 4 kHz (SLF) and 22 kHz (PHF) raw signals were recorded permanently. Due to strong interference with the EM710 multibeam bathymetry profiler, the pulse interval was selected by the operator between 500 ms and 2000 ms (single pulse mode) depending on the anticipated slope angles and the amount of interference observed at a given water depth. The water velocity was set to 1500 ms-1. Technical problems occurred rarely and could be solved during the cruise. The overall data quality and coverage was good. All raw data (for PLF, PHF, SLF and SHF frequencies) were stored in the ASD data format (Atlas Hydrographic), which contains the data of the full water column of each ping as well as the full set of system parameters. Additionally, a 200 m-long reception window starting at 90% of the seafloor depth was recorded in SEG-Y and compressed PS3 data formats after resampling the signal back at 12.1 kHz and converting from trace amplitude to envelope. This format is in wide usage in the PARASOUND user community and the limited reception window provides a detailed view of sub-bottom structures. All data were converted to SEG-Y format during the cruise using the software package ps32sgy (Hanno Keil, Uni Bremen). The software allows generation of one SEG-Y file for longer time periods, frequency filtering (low cut 2 kHz, high cut 6 kHz, 2 iterations), and subtraction of mean. All data were loaded to the seismic interpretation software IHS Kingdom. The entire PARASOUND data set will be transferred to

international data banks and may be used by specialists for further shore-based processing and analyses.

#### **5.1.2 Preliminary Results Bathymetry and Subbottom Profiling**

a) Multibeam echosounder

A total of 3340 km2 of multibeam echosounder data were acquired during M191 (Figure 5.3). The data are generally of good quality, although we did observe an interference between the EM710 and Parasound, especially at shallow depths, which resulted in significant noise in some of the bathymetric and backscatter data. The most significant observations made include new volcanic centres: (i) between Galatea and Anfitrite volcanoes (Figure 5.4), (ii) NE of Linosa island (Figure 5.5) and (iii) Nameless Bank (Figure 5.6a). The backscatter signatures of these volcanic centres were of very good quality (Figure 5.6b), and was used to guide the dredging operations. Flares in the water column data, indicative of active venting/seepage, were observed at a number of locations (Figure 5.7).





**Fig. 5.3:** Spatial coverage of multibeam echosounder bathymetry data acquired during M191.



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Fig. 5.4: Multibeam echosounder 
  bathymetry data of new 
  volcanic center located 
  between Galatea and 
  Anfitrite volcanoes.
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**Fig. 5.5:** Multibeam echosounder bathymetry data of a new volcanic center located NE of Linosa.



Fig. 5.6: a) Multibeam echosounder bathymetry (upper panel) and (b) backscatter data of a new volcanic centre located on Nameless Bank (lower panel).



Fig. 5.7: EM710 water column data image. Flare (green elongated feature in the centre) suggesting active venting or seepage.

b) Sediment echo-sounding system

The PARASOUND P70 shows overall good penetration into the subsurface, except for areas where coarse-grained sediments, bedrock, or steep slopes scatter the transmitted energy and distort the proper imaging of the subsurface. Overall 2490 km PARASOUND profiles have been acquired during M191 (Figure 5.8). These data provide useful information to identify exposed and buried volcanic centres (Figure 5.9) and the adjacent stratigraphy, which can be used to deduce a relative age of volcano formation. In addition, sub-bottom profiles adjacent to volcanic centres frequently showed evidence of mass transport deposits (Figure 5.10) and fluid flow structures (Figure 5.11).





Fig. 5.9: Sub-bottom profile across Galatea volcano.



**Fig. 5.10:** Sub-bottom profile across a mass transport deposit west of Nameless Bank.



Fig. 5.11: Sub-bottom profile across pockets of gas just south of Sicily

## **5.2 Magnetic Field Survey**

## (Jonathan Ford, G. Matilde Ferrante)

Marine magnetic surveys involve measuring the Earth's magnetic field by towing a magnetometer behind a ship. Lateral magnetic susceptibility contrasts generate sharp anomalies (peaks or troughs) in the otherwise smoothly varying magnetic field. This means that the method is sensitive to the presence of contacts between sedimentary and magmatic bodies, which have very different magnetic susceptibilities due to the relative abundance of magnetic minerals in igneous rocks. The method can therefore provide an indication of the composition of submarine structures identified in the bathymetric data. For the SUAVE cruise, the magnetic anomaly data is primarily used to discriminate between volcanic cones and non-volcanic seamounts (e.g., carbonate mounds).

## **5.2.1 System Overview and Data Processing**



Fig. 5.12 (above): Towfish containing Overhauser magnetometer.

**Fig. 5.13** (right): Cable on reel (300 m).



The magnetic data were acquired with an Overhauser proton procession magnetometer, which contains a chamber filled with a liquid rich in hydrogen. Electrons dissolved in the liquid are excited by a radio frequency power source and pass on their energy to the protons, altering their spin states. The transfer of energy from electrons to the protons in the hydrogen atoms is called the Overhauser Effect. Once the protons are spinning, the power is removed, and the protons spiral back to their original alignment with the external geomagnetic field. The rate of precession is dependent on the total intensity of the magnetic field and is measured.

The system used during Exp. 191 consists of:

- 1) A "SeaSpy" towfish unit (Marine Magnetics Corporation) that contains the magnetometer and driving electronics (Fig. 5.12);
- 2) An isolation transceiver for powering and communicating with the towfish;
- 3) A high-strength marine tow cable containing a single twisted wire pair (total length 300 m), housed on a reel (Fig. 5.13);
- 4) A deck leader cable to connect the cable reel to the isolation transceiver;
- 5) An RS232 interface cable that connects to isolation transceiver to a standard PC RS232 port;
- 6) A universal input power supply that powers the isolation transceiver and provides a 48V power supply to the towfish.





Measurement of the magnetic field is done completely inside the towed fish. The tow cable supplies power to the towfish and provides a bidirectional digital communication link (Fig. 5.14). The system has a maximum resolution of 0.001 nT and a maximum sample rate of 4 Hz. The software used for the acquisition and real-time visualisation and quality control is called "SeaLink" (Marine Magnetics Corporation), which runs on a Windows 10 operating system.

Magnetic surveys traditionally have several parallel lines perpendicular to the expected orientation of the anomalies, in order to make a grid. During M191 the magnetic data were acquired opportunistically, with the intention of using the in-profile anomalies to confirm the composition of sea mounts. Due to this, sail lines instead followed the multi-beam echosounder mapping survey, resulting in magnetic profiles which have variable orientation. The magnetometer sample rate was 0.3 Hz, except for areas where rapid changes in magnetic field and/or the presence of very short wavelengths anomalies (e.g., the Tetide survey area, in around 50 m water depth) required a higher sample rate of 1 Hz.

The deployment procedure was as follows:

- 1) Establish communication with the tow fish on the deck;
- 2) Synchronize the towfish clock with the GPS time;
- 3) Pay out the whole reel of cable (except for a couple of rounds left on the reel) assisted by a deck winch and a rotating drum on the edge of the ship;

4) Secure the cable to the ship with a fixed line, remove the tension in the cable between the reel and the tow point.

The recovery procedure was as follows:

- 1) Slow ship to 4 kn, wind cable back onto reel with assistance from deck winch.
- 2) When towfish is on deck, immediately (whilst towfish internal temperature is still at operating temperature in water) perform depth test to check that the pressure (depth) sensor reads 0 m, to ensure that the zero-calibration correct.

Both deployment and recovery took approximately 10-15 minutes.

The magnetic survey measured the total magnetic field intensity (in nT) with the magnetic anomaly being obtained from the total field intensity by subtracting:

- 1) The International Geomagnetic Reference Field (IGRF), which accounts for the magnetic field generated by the Earth's core;
- 2) The diurnal effects generated by the variations of the external magnetic field (largely from space weather). This was measured by the Lampedusa Magnetic Observatory (INGV), with a sample interval of 1 minute. Data for the previous day were made available to download from 00:00 UTC.

In order to provide real-time assessment of the magnetic bodies before dredging operations, preliminary grids were built onboard (see 5.2.2 below). No between-profile levelling was performed at this stage.

## **5.2.2 Preliminary Results Magnetic Surveys**

In the following, exemplary results of specific magnetic survey areas are shown (Fig. 5.15, 5.16) and the coordinates of the conducted survey lines (start and end points) are provided (Fig. 5.17, Table 5.1).





Fig. 5.17: Geographical overview of the magnetic surveys conducted during Exp. M191 in the Sicilian Channel.





## **5.3 Dredging**

**(**Jörg Geldmacher, Sebastian Watt, Alastair Hodgetts, Folkmar Hauff, Maxim Portnyagin, Aaron Micallef)

#### **5.3.1 Methods, Shipboard Procedure and Shore-based Analyses**

## (J. Geldmacher)

The search for potential dredge sites was guided by 1) predicted bathymetry, derived from gravity data and ship depth soundings (EMODnet Bathymetry: https://emodnet.ec.europa.eu/en/bathymetry), 2) our own multibeam mapping (see Chapter 5.1) and 3), the results of the magnetometer surveys (if available) (see Chapter 5.2). In addition, existing multi-beam data were available for limited parts of the working area, which were obtained from published sources or provided by our Italian colleagues (Emanuele Lodolo, Dario Civile (OGS), Danilo Cavallaro (INGV). Final positioning of the vessel at each dredge station also included considerations of wind, swell and drift conditions. During M191, however, favourable weather conditions allowed maximum flexibility in the selection of dredge track directions, except for the short time span from July 26.-27., when wind speed and wave height transiently increased. Dredge tracks were usually located - depending on the morphology of the structures - on steep slopes of scarps, canyon walls, fault zones, and the flanks of cones, ridges, and larger seamounts to avoid thick sediment cover. The small volcanic structures in the Sicilian Channel, however, often possess very shallow slopes, reducing the likelihood of a direct contact of the dredge with exposed hard rock.

A considerable challenge was the dense distribution of submarine cables across the entire working area. A careful protocol was followed to ensure that all potential dredge sites were first assessed based on cable location, and no sites were targeted where the location was within 1 km and <2.5 times the water depth from the nearest cable, based on whichever value was greater. As a result, several locations: Doride (Lodolo et al., 2019); various cones along the Graham Bank (Cavallaro and Coltelli, 2019); Tetide (Calanchi et al., 1989); and the majority of the Linosa III seamount structure could not be dredged. In addition, a region around Pantelleria was not permitted for access, and a marine protection zone around Linosa (as well as cable positions) limited access to the cones north-west of Linosa to those at the far northern limit of the submerged volcanic platform (cf. Romagnoli et al., 2020).

A final challenge to the recovery of volcanic material was the accumulation of unconsolidated seafloor sediment, especially in the northernmost sites of the study area and close to terrigenous sediment sources from Sicily, and in areas where a significant build-up of biogenic carbonate on the seafloor occurred, either in the form of rhodolith gravels or coralline and other bioclastic accretions. The latter was the case across much of the study area. In an effort to minimize recovery of carbonate material, seafloor backscatter data were used, targeting sites with high backscatter wherever possible. Nevertheless, many dredges recovered only marine surficial carbonate, even in sites of confirmed volcanism (e.g., the Graham Bank cones; Cavallaro and Coltelli, 2019).

Rock sampling on cruise M191 was carried out using rectangular chain bag dredges (Fig. 5.18). Chain bag dredges are essentially large buckets with a chain bag attached to their bottom and steel teeth at their openings, which are dragged along the ideally sediment-free ocean floor by the ship's winch.



**Fig. 5.18:** Deployment of a chain bag dredge near the Italian island of Linosa.

If volcanic rocks (or other rocks which appeared worthwhile sampling) were obtained, they were sorted and selected for further processing. First, these were cleaned and cut using a rock saw. They were then examined with a hand lens and binocular microscope, and grouped according to their lithologies and degree of submarine weathering. The immediate aim was to determine whether material suitable for geochemistry and radiometric age dating had been recovered. Best suitable volcanic samples have an unaltered groundmass, empty vesicles, glassy rims (ideally), and -if applicable- well-preserved phenocrysts. If suitable samples are present, the ship moved to the next station. If they were not, then the importance of obtaining samples from the respective site was weighed against the required time commitment for repeating a dredge haul. Due to the generally shallow water depth in the Sicilian Channel (resulting in relatively quick individual dredge hauls), unsuccessful dredges on the same target were frequently repeated.

Fresh blocks of representative (igneous) samples were then cut for post-cruise thin section and microprobe preparation, geochemistry and further procedures, to remove manganese and alteration products, and/or to extract volcanic glass (if present). Each of these sub-samples, together with any remaining bulk sample, was described, labelled, photographed, and finally sealed in plastic bags for transportation to GEOMAR.

Igneous rocks sampled during M191 from the ocean floor will be analysed using a variety of different geochemical methods: Ages of suitable rock samples will be determined by <sup>40</sup>Ar/<sup>39</sup>Ar

laser step-heating dating. Major element geochemistry by X-ray fluorescence (XRF) and electron microprobe (EMP) will constrain magma chamber processes. Trace element data, obtained by inductively coupled plasma mass spectrometry (ICP-MS), will help to define the degree of mantle melting and help to characterize the chemical composition of the source. Phenocryst assemblages and compositions will be used to quantify magma evolution. Petrologic studies of the volcanic rocks will also help to constrain the conditions under which the melts crystallized. The composition of mafic basalts and basaltic glasses, as well as mafic melt inclusions, can be used to assess mantle temperatures at which melting took place, as well as pressures and degrees of melting. Sr, Nd, Hf and Pb (double spike) isotope ratios, determined by thermal ionization mass spectrometry (TIMS) and multi-collector ICP-MS, reflect the long-term evolution of the magma source(s) and thus serve as tracers to identify mantle domains and possible crustal contamination (e.g., from the surrounding carbonate platform that all magmas have to pass through). Morphological and volcanological studies will constrain eruption processes, eruption environment and evolution of the volcanic structures

Representative non-magmatic rocks i.e., carbonate crusts or limestone samples from the carbonate platform were also collected and can be transferred to co-operating specialists for further shore-based analyses.

#### **5.3.2 Preliminary Results Dredging**

(Sebastian Watt, Alastair Hodgetts, Folkmar Hauff, Maxim Portnyagin, Aaron Micallef)

Sixty-eight sites, selected across the study area, were dredged in total. Twenty-four of these sites recovered volcanic material (Table 5.2; Figure 5.19). Only five dredges were empty, with the remainder retrieving carbonate material or mud/sediment. Most of this carbonate material, a constituent in nearly all dredges, was young surficial biological material of various types. Ten sites, however, retrieved denser carbonate in the form of broken blocks or pebbles, inferred to be derived from the local bedrock. Dredge sites were chosen based on a combination of indicators:

1. Reference in existing scientific literature to volcanic centres identified either by prior sampling of volcanic rock (e.g., Beccaluva et al., 1981; Calanchi et al., 1989; Rotolo et al., 2006), prior seafloor mapping of magnetic anomalies (Lodolo et al., 2012, 2019; Civile et al., 2018), or prior interpretation of volcanic landforms (Calanchi et al., 1989; Coltelli et al., 2016; Cavallaro and Coltelli, 2019). The site of the 1831 eruption of Ferdinandea/Graham Island (Coltelli et al., 2016) was also targeted, and more uncertain reports of other historical events (Washington, 1909; Francis, 1995; Siebert and Simkin, 2023) were revisited to locate potential eruption sites.

2. Interpretation during the expedition of our newly acquired bathymetric and totalfield magnetic datasets, in order to identify magnetic anomalies, particularly those that corresponded to positive relief seafloor landforms.



**Figure 5.19:** Map of of the main working region showing all dredging sites, distinguishing those that recovered volcanic material.

In several instances, our approach of using pilot seafloor mapping to evaluate the magnetic signature of bathymetric highs demonstrated that structures present on low-resolution GEBCO/EMODnet datasets were not present within our new multibeam bathymetry data (and did not display magnetic anomalies). In other instances, previously postulated but unsampled volcanic seamounts (e.g., Linosa I and Linosa II; Calanchi et al., 1989) were found to have no magnetic signatures. Accordingly, dredging targets were narrowed down to a small number of sites where we had the greatest confidence of potential volcanic seafloor outcroppings. There were some instances where magnetic signatures were found to only roughly correspond to bathymetric heights (e.g., Cimotoe; Calanchi et al., 1989; Terribile Bank). In these areas, magnetic signatures likely represent subsurface or partially buried igneous rock, and in these locations, we targeted sites with positive and steep topography that were deemed to provide the best opportunity of reaching the underlying bedrock.

In the following, we divide the study area into sub-regions (Table 5.2; Figure 5.19). These were broadly targeted in the sequence described, although some areas (Terribile Bank; Nameless Bank) were returned to at later points in the expedition.

Sub-region	<b>Dredge</b>	Summary of volcanic material recovered
	stations	
N end of Capo Granitola	5, 6, 7, 8,	1. laneira cone: one dense lava fragment (site 7) and block with
fault	10, 11,	mafic volcaniclastic fragments in altered matrix (site 8)
(for cone locations refer	12, 13,	Nesea cone: one glassy dark vesiculated clast (site 13) 2.
to Lodolo et al., 2019)	14	
Graham Bank	16, 17,	1. S6 cone: blocks with scoriaceous volcaniclastic fragments in
(for cone locations refer	20, 21,	altered matrix (sites 16, 17)
to Cavallaro and Coltelli,	22, 23,	S5 cone: hydrothermal chimney structures (site 20) and 2.
2019)	24, 34,	dense mafic lava blocks (site 21)
	35, 36	3. S4 cone (Ferdinandea): mixed volcaniclastic material (scoria,
		inflated blocks, palagonitised tuff fragments, dense blocks) (site 23)
		4. S1 cone: block with scoriaceous volcaniclastic fragments in
		altered matrix (site 24)
		5. Site of S10 cone: coarse ash/very fine lapilli, sparse mafic
		fragments embedded in surface mud (may not be locally derived) (site
Adventure Plateau	26, 28,	34) None
(Galatea, Anfitrite, Tetide;	29, 30,	
cf. Calanchi et al., 1989)	31	
<b>Terribile Bank</b>	37, 38,	1. E Terribile bank: Ignimbritic surficial block, with brown fine-
	39, 40,	grained matrix encasing pumice clasts, crystals and carbonate lithics
	41, 42,	(site 37)
	43, 44,	N Terribile bank: Green rounded blocks of clastic 2.
	45, 68,	hydrothermally derived/altered material associated with pockmark
	69, 70,	field (site 70); multiple blocks of pervasively altered crystalline rock
	71, 72,	from edge of pockmark (site 73)
	73, 78,	
	79	
Nameless Bank	47, 48,	Upper cliff and associated slopes on N side of ridge projecting 1.
	49, 75,	on far E side of Nameless Bank, retrieved across distance of several
	76, 81,	km: dense (non-vesicular) porphyritic mafic lava blocks, variably
	82, 83,	altered and with slight variation in phenocryst assemblage (sites 83,
	84, 85,	84, 85, 87, 88), alongside flattened red clasts of altered fragmented
	87, 88,	volcanic material. Uppermost site, above cliff, contained minor coarse
	89, 90	volcanic clasts embedded in poorly lithified carbonate sand pebbles
		(site 89).
Linosa and nearby	51, 52,	Cone at far N edge of platform NW of Linosa: Vesicular mafic 1.
seamounts	54, 55,	lava blocks (sites 54 and 55), including peperitic texture
	56, 57, 59, 60,	2. Adjacent cone immediately SE: Vesicular mafic lava blocks (site 56)
	62, 63,	Linosa III, far E edge: small pebbles of calcite-cemented 3.
	64, 65,	scoriaceous clasts embedded within lithified carbonate mud (sites 60
	66	and $64)$

**Table 5.2: Summary of dredge sites by sub-region**

## N end of Capo Granitola fault

This region preserves some of the youngest volcanic landforms in the study area (Figure 5.19), in the form of isolated small cones or rings with strong magnetic anomalies (Lodolo et al., 2019). Sub-bottom profiles indicate that these structures date from around the Last Glacial Maximum, with the bases of each structure being partially buried by younger sediment (Lodolo et al., 2019). The largest, western-most center of Actea is the most degraded landform. Doride is a well-shaped truncated cone but was inaccessible to dredging because of nearby submarine cables. The shallow setting of these sites, in water depths of just 40 m for the northern cones, presented a challenge for dredging. Their setting, close to the Sicilian shoreline, shallow depth also likely leads to a relatively high sedimentation rate and, despite their young age, has ensured a sufficiently thick biogenic/carbonate drape which makes it difficult to retrieve material from the underlying volcanic structures.

Actea was dredged twice (sites 5 and 6) recovering only fragments of encrusted bioclastic carbonate. Ianeira, in slightly deeper water to the south, was dredged at sites 7 and 8, the first recovering a single piece of moderately vesicular mafic lava amongst multiple bioclastic carbonate fragments, and the second retrieving fragments of palagonitised tuff with dispersed grey angular to sub-rounded mafic lapilli, among several other large coral-encrusted carbonate blocks. A nearby cone, Ianassa, was dredged twice (sites 10 and 11) without recovering any volcanic material, collecting only a variety of coral fragments. Nesea, further north and in shallower water (70 m), has a low-profile concentric morphology, suggestive of a tuff ring. Two dredges (sites 12 and 13) recovered large volumes of mud with coral and shell fragments, and one volcanic rock at site 13, a glassy dark mafic pyroclast with an open foamy texture. The final volcanic center in the group, Climene, has a similar morphology to Nesea but is more degraded. A dredge there (site 14) recovered only mud with several large coral fragments and bivalves.

The volcanic rocks retrieved from this region are the first samples recovered from these volcanoes, except for a coral-encrusted lapilli tuff from Actea that was dated in Lodolo et al. (2021).

#### Graham Bank

The Graham Volcanic Field cones comprise a N-S orientated group of ten cones (Cavallaro and Coltelli, 2019), interpreted as being generated by monogenetic explosive eruptions in a shallowsubmarine to emergent environment. This interpretation is supported by observations made during the 1831 eruption of Ferdinandea (cone S4; Cavallaro and Coltelli, 2019), which briefly formed an island before erosion led to its current form of a truncated submarine cone with a wide wavecut platform. Other cones in the field have a similar morphology, with more deeply submerged centers having narrow summit platforms and the deepest having a simple conical morphology. Unlike the dispersed monogenetic cones further north, the Graham Volcanic Field includes three clusters of closely spaced cones, suggesting reactivation of feeder structures, or events that were closely spaced in time within each cluster.

From the southernmost cluster of three cones, only S1 (Cavallaro and Coltelli, 2019) was accessible for dredging during M191, due to submarine cable locations. Here, a single dredge (site 24) recovered one volcaniclastic block, preserving angular mafic scoriaceous clasts within a green and white altered matrix, alongside several large rounded carbonated blocks. Further north lie the paired cones of Ferdinandea (S4) and the morphologically similar but much larger S5. A single dredge on Ferdinandea (site 23) recovered a large volume of loose pyroclastic debris, likely representing the wave eroded material derived from the former island and present-day wave cut platform, accumulated on the outer surface of the submerged cone. A mix of clast types (Figures 5.20 and 5.21) included breadcrusted inflated blocks, dense vesiculated fragments, rounded palagonitised tuff gravel, scoriaceous lapilli and two sedimentary xenoliths. On the nearby S5 cones, a dredge site on the NE side (site 20) retrieved several pieces of hydrothermal chimney constructions, comprising cemented sand-sized volcanic clasts with a central pipe (Figure 5.21Bi, ia). These were recovered alongside altered material and bioclastic carbonates. They are likely associated with an active vent field imaged in water column data, on the S side of a large pockmark, and represent fumarole structures. On the SW side of the same cone (site 21), fresh volcanic rock was recovered, in the form of dense, poorly-vesicular clasts, alongside a large volume of bioclastic carbonate. A dredge site at the base of the cone (site 22, west side of the cone), across a rough area of seafloor that may represent a debris field or lava flow, recovered no material.

Two isolated cones lie further north, but only one of these (S6) could be dredged due to submarine cable positions. Dredge sites 16 and 17 recovered mostly bioclastic carbonate and coral fragments, with some palagonitised tuff blocks preserving dark, angular and finely vesicular clasts. The northernmost cluster in the Graham Volcanic Field comprises a number of small cones that were again mostly inaccessible due to cable positions. The western cone was dredged (sites 34- 36) but both dredge hauls recovered mud and a small number of bioclastic carbonate fragments. The only volcanic material comprised sparse coarse mafic ash and lapilli  $(\leq 0.8 \text{ cm})$  grains accumulated in grey surface clay, embedded in some of the carbonate blocks recovered from haul 34. We cannot be confident that these are derived from this cone, and they may represent fallout deposits from another eruption in the area, such as the 1831 Ferdinandea eruption.

#### Adventure Plateau

South of the Graham Volcanic Field, a volcanic centre named Cimotoe is marked in several papers (Calanchi et al., 1989; Civile et al., 2018), although with slightly ambiguous positioning. A cluster of five conical structures visible in EMODnet bathymetry and aligned with the Graham Volcanic Field was initially mapped, but several of these structures were found not to exist when the area was mapped with high-resolution multibeam bathymetry during M191, and there is no distinct magnetic anomaly in the area that corresponds to a positive seafloor landform. A track to the NW, crossing onto Adventure Plateau, was then mapped, transiting over small positive landforms, which were found not to have a clear magnetic anomaly. A dredge taken on one of these (site 26) recovered only bioclastic carbonate blocks.

Further NW, an alignment of several large but highly eroded volcanic landforms has been previously identified, comprising the three centers of Tetide, Anfitrite and Galatea (Calanchi et al., 1989), with the first two being previously sampled and demonstrated to be volcanic. These are far larger than the Graham Bank and Capo Granitola cones, and likely represent much longerlived, polygenetic structures. Their high level of degradation suggests a much older age, and these landforms would have been exposed above sea level at the last glacial maximum (cf. Civile et al., 2015). Tetide could not be dredged due to the location of submarine cables. Anfitrite was dredged on the eastern flank, but this recovered only several surficial carbonate blocks, with an open encrusted bioclastic structure (site 28), and nearby sites on a newly identified volcanic landform between Anfitrite and Galatea (29 and 30) also only recovered carbonate (dense bioclastic) material. Galatea was located based both on bathymetry (Fig. 5.3) and its corresponding magnetic anomaly, with a single dredge on its eastern slope (site 31). Again, this recovered only dense carbonate blocks.







#### Terribile Bank

Terribile Bank has been cited as containing volcanic landforms in several papers (Falzone et al., 2009; Civile et al., 2018; Cavallaro and Coltelli, 2019). Civile et al. (2018) provide one sub-bottom and magnetic-anomaly profile that suggests the presence of magmatic bodies, but otherwise the interpretation of volcanic centers is based solely on the varied surficial landforms in the area, some of which have irregular, lava like morphologies or represent small rounded or truncated conical forms. Our initial dredges targeted areas identified in Civile et al. (2018) and Cavallaro and Coltelli (2019) as volcanic landforms, while a later dredge visited the MAC-06 site of Falzone et al. (2009). These sites (37-41, 45) recovered only carbonate material, including many sites with rhodoliths, varied coralline material and bioclastic carbonate. At site 68, the rounded form interpreted as a small volcano (MAC-06) by Falzone et al. (2009), the dredge recovered lithified beach or coastal carbonate sand, alongside dense carbonate pebbles, with further varied carbonate lithologies recovered at site 69, across the same structure, suggesting that it is not volcanic. The only other dredge haul from this part of Terribile Bank that recovered material interpreted as bedrock carbonate (as opposed to surficial lithified bioclastic mud, rhodoliths, or biogenic concretions) was conducted at site 37, at the eastern edge of a lobate mounded structure. This retrieved varied carbonate pebbles and was also the only site in the immediate area to recover non-carbonate material, in the form of blocks with a brown fine-grained matrix with dispersed carbonate clasts and crystals, interpreted as peperitic (Fig. 5.21Ai,ii).

All of the above sites at Terribile Bank did not show a close association between magnetic anomalies and surface landforms, despite the prior interpretation of young volcanic structures in the area. However, a number of strong magnetic anomalies were mapped on Terribile Bank, suggesting local igneous bodies and consistent with the observations of Civile et al. (2018). One prominent anomaly is on the north side of the bank. This was targeted at dredge sites 42-44 (targeting prominent topographic highs and areas with high backscatter, but again retrieving only surficial carbonate, including rhodoliths), and then later at sites 70-73. These sites were more closely associated with a magnetic anomaly, much of which did not overlap with surficial topography. In this area, on the southern edge of the anomaly, a fault bound valley with intensive pockmarks lies between two topographic highs. Sites 71 and 72, on the highs, recovered only surficial carbonate, but 70 and 73 retrieved highly altered material consistent with hydrothermal venting. At site 70, on the western edge of a pockmark on the W side of the valley, the sample contained a dense bedrock carbonate block alongside green rounded and soft blocks, containing crystals and secondary minerals, potentially derived from an extremely altered igneous protolith. Such an interpretation is supported by the range of samples at site 74, on the margin of an adjacent pockmark. This sample contained no carbonate material, but numerous dense and crystalline volcanic, porphyritic blocks, several with angular fracture planes and all displaying varying degrees of pervasive alteration, in most cases entirely replacing primary phenocryst minerals. These observations are consistent with a shallow subsurface igneous body being the source of the magnetic anomaly in the area, rather than young volcanic constructions.

Further magnetic mapping revealed additional magnetic anomalies in various areas on Terribile Bank. Again, these generally do not correspond closely to surface landforms, but one exception is a strong anomaly in the eastern central part of the bank, where a linear anomaly corresponds to a topographic ridge. This ridge has surface morphologies consistent with carbonate areas on the bank, but given its close association with a magnetic anomaly, it was dredged at sites 78 and 79. This did not recover any volcanic material, but did retrieve rounded carbonate pebbles and a cemented carbonate sand, suggesting subaerial exposure and potential carbonate constructions on top of subsurface igneous rock.

#### Nameless Bank

Nameless Bank has previously been dredged three times, with igneous rocks described in Beccaluva et al. (1981), Calanchi et al. (1989) and Rotolo et al. (2006). All these dredge sites are located on the N side of a ridge that extends from the eastern side of Nameless Bank (Fig. 5.6a). Three initial dredges (sites 47-49), targeting the same areas, did not recover any volcanic material. The first two recovered only clay, while the third retrieved varied carbonate lithologies, including breccias and cherty material, interpreted as representing bedrock exposed in the north wall of the ridge. Further mapping confirmed the presence of strong magnetic anomalies across the ridge, and additional dredges targeted areas of high backscatter (sites 75 and 76) but again recovered only dense carbonate (mid part of the ridge, site 75) or reef coralline and shelly material (lower ridge, site 76). Further dredges targeted sites further east along the cliff (sites 81 and 82) and then a high backscatter area below the upper cliff (site 83), finally recovering numerous volcanic fragments, mostly highly altered and most frequently in the form of angular flattened and reddened clasts, with a fragmental fabric. Subsequent dredges targeting the cliff above this slope, and similar levels along the upper parts of Nameless Bank, retrieved blocks of dense volcanic rock with fresh broken surfaces (dredge sites 84, 85, 87, 88). These were all very dense, non-vesicular porphyritic blocks with an aphanitic and in some cases likely glassy matrix. Phenocryst assemblages vary slightly across the sites, which are spread over several kilometres, but are generally highly altered. Brown altered olivine is abundant in most specimens, rarely with fresh material in the centre of crystals. Black, tabular pyroxenes are generally less altered and of lower abundance, but still frequent. Some samples had large plagioclase laths, but others had no visible phenocryst plagioclase. The sampled sites correspond well to the intense magnetic anomaly on Nameless Bank and suggest that the upper parts at least are of volcanic construction, with in-situ outcrops of large, eroded lava bodies. The top of Nameless Bank shows a terraced but smooth morphology, likely extensively draped by sediment, with sparse outcrops in the centre. The margins of the terraces retrieved coarse subrounded pebbles of bioclastic sandstone, in some cases with 5-10% volcanic grains (site 89).

#### Linosa and nearby cones

Several bathymetric highs between Nameless Bank and Linosa were transected in an initial profile in order to verify if these corresponded to magnetic anomalies. The previously postulated seamounts of Linosa I and II (Calanchi et al., 1989) show no magnetic anomaly. In contrast, a strong and complex anomaly was observed over Linosa III and, as expected, on the submerged flanks of Linosa itself.

Dredging was conducted offshore Linosa in the only accessible region (given cable locations and a marine protection zone), which was at the far end of the submarine volcanic ridge NW of Linosa. Here, two separate cones were sampled. The first dredges at the most northerly cone at the end of the ridge (sites 51 and 52) recovered no material except for mud and small carbonate fragments (a later dredge here, site 57, again recovered only mud), but subsequent dredges recovered vesicular, mafic, porphyritic lava blocks and fluidal fragments (sites 54 and 55). A further dredge (site 56) targeted the lower flanks of the cone immediately to the SE, retrieving dense and finely-vesicular porphyritic lava blocks.

Mapping at Linosa III revealed rounded circular landforms closely corresponding to a magnetic anomaly, interpreted as volcanic. Cable locations meant that only sub-optimal sites, at the far eastern edge of these structures and on a faulted boundary, could be targeted. Site 59 recovered dense carbonate brecciated blocks, interpreted as bedrock from the fault scarp, alongside other carbonate lithologies. Site 60, slightly upslope and closer to the volcanic landforms, retrieved a bioclastic carbonate mud encasing a rounded, flattened pebble or carbonate-cemented scoria fragments, interpreted as being locally derived and then set within younger lithified bioclastic carbonate on the current seafloor. A very similar specimen was recovered in a dredge nearby (site 64), but other dredges in the vicinity were either empty (site 62 and 66) or recovered only bioclastic carbonate and encrustations (sites 63 and 65).

#### Biogenic and carbonate materials recovered at various locations

The dredging collected numerous and varied carbonate samples from across the study area (Fig. 3.2). This included carbonate bedrock in places, but for the most part comprised surficial and recent material, giving insights into conditions and biota on the present-day seafloor. This sample set holds potential significance in evaluating facies in bathymetric and backscatter data, including constraints on specific environments such as the presence of rhodolith gravels. The encrusting organisms and other material, particularly the extensive range of coral fragments observed at a number of sites, may also give new insights into the nature and range of seafloor colonisation.
# **5.4 Nannoplankton Sampling**

(Odysseas A. Archontikis)

# **5.4.1 Introduction and Methods**

Coccolithophores (Prymnesiophyceae, Haptophyta) are unicellular marine phytoplanktonic organisms distributed worldwide and one of the most prominent groups of primary producers with a key role in sediment formation. Through photosynthesis and calcification, coccolithophores play a major role in regulating the marine food web and are a key buffer of the ocean carbonate chemistry. As a result, they have attracted an increasing and interdisciplinary interest with studies of their biogeography, biodiversity and calcification serving as means for monitoring plankton response to climate change. This interest has been matched for the Mediterranean Sea, with few large-scale surveys of coccolithophores conducted in the area (Knapperbutsch, 1993; Kleijne, 1993; Oviedo et al. 2015; D' Amario et al. 2017) and new species of them having been discovered and described (e.g., Cros & Fortuño, 2002; Archontikis et al. 2020; Archontikis & Young 2021; Archontikis et al. 2023a; Archontikis et al. 2023b).

M191 Expedition provided a unique platform to perform nannoplankton sampling across the Mediterranean Sea (from Algeciras, Spain to Piraeus, Greece) and therefore, investigate living plankton assemblages and determine patterns of coccolithophore/nannoplankton biogeography, calcification, and life-cycle dynamics. Accordingly, surface seawater samples were collected in the main Exp. M191 working area in the Sicilian Channel and during transits throughout the Spanish and Greek exclusive economic zones, to assess extant coccolithophore (nannoplankton) species assemblages. Surface water samples  $(\sim 12 \text{ L})$  were collected day and night at approximately every 2–4 hours whilst in transit and approximately every 12 hours while dredging operations were carried out. Sampling was conducted either via a plastic bucket with a nylon rope over the side of the starboard main deck or via a tap near the suction valve of the automated thermosalinograph ("Reinseewasser", supplied by the sinus pump system) located at the bow of the vessel. The sampling depth approximately represents a mixed upper 10 m of surface water. Upon retrieval, the plastic 12L bucket was transferred into the lab; latitude and longitude coordinates, the time and date of each collection (in UTC) and the water temperature, salinity, conductivity, and density were recorded (see Table in APPENDIX 11.2).

For each sampling station, c. 5–8 L of seawater were filtered using a stainless-steel filtration ramp onto 47 mm diameter, 0.8 µm pore-size Millipore polycarbonate track-etched filters (Fig. 5.23). Prefiltration through a brass 63 µm test sieve was carried out to remove larger zooplankton and contaminants. A low-pressure vacuum pump was also used during the filtration to avoid mechanical breakage of specimens. Salt was removed by washing the filters with ~3 mL of Milli-Q deionised water buffered to a pH of about 8.5 with ammonia (NH3). After filtration, the filters were placed individually in 47 mm plastic Millipore Petri-slides and dried in an oven at 50°C. Once dry, a portion of each filter was cut out and mounted onto a glass microscope slide using Norland Optical Adhesive (NOA) 61 (refractive index= 1.56). The filter membranes were then sealed in Petri-slides, covered with Parafilm, and frozen for post-cruise analyses.





## **5.4.2 Preliminary Results of Nannoplankton Sampling**

Eighty-two water samples were collected during the passage across the Mediterranean Sea from Algeciras, Spain to Piraeus, Greece and in the main working area in the Sicilian Channel where dredging operations were conducted. Water temperature, salinity, conductivity, and density were measured/recorded by the ship's built-in automated thermosalinograph at all samplings locations, and these fluctuated, respectively, from 19–31°C, 36–38 PSU, 50–61 mS/cm and 1023– 1025 kg/m<sup>3</sup>. The bulk of the nannoplankton analyses will be undertaken post-cruise, together with synthesis of data from previous coccolithophore studies. Some preliminary observations are, however, made by using a Leica/Wild M3Z Stereo Microscope. These revealed the presence of 1) organic matter ranging in content across sampling stations; 2) larger planktonic organisms (e.g., juvenile planktonic foraminifera) that were found sporadically on the filter membranes; and 3) diverse assemblages of smaller  $(< 45 \mu m$ ) plankton. Detailed nannoplankton analyses including taxonomic identification, counts and geochemical studies will be implemented post-cruise with appropriate access to geochemical labs and high-resolution microscopy (partly in cooperation with J.O. Herrle, Goethe University, Frankfurt).

In total, the 82 water samples collected during M191 expedition, represented over 510 L, while preliminary analyses were carried out on ship on 28 samples. Seawater sampling was productive in terms of recovery of plankton assemblages and will correspondingly formulate promising post-cruise research on coccolithophore diversity, ecology and calcification as means to understand and monitor effects of climate change on marine ecosystems.

# **6 Ship's Meteorological Station**

# (A. Raeke)

On the morning of 16/07/2023 at around 10 am, the research vessel RV METEOR left the port of Algeciras, Spain. A low-pressure system centered over the east coast of Spain was slowly moving southwest and weakening. Otherwise, low-pressure contrasts prevailed over much of the Mediterranean. The initially light winds in the Strait of Gibraltar later increased to westerly 6 Bft (jet effect). From midday, the wind decreased to 2-4 Bft from various directions, with significant wave heights of 1 m and low swells, which also decreased. With continued low air pressure contrasts in the Mediterranean on the following days, the wind on 17/07/2023 was still fresh with 5 to 6 Bft from northeasterly directions with a significant wave height of 1,5 m.

On 18/07/2023, the wind veered southeast and decreased to 2-3 Bft before temporarily freshening to west 4 Bft in the evening of 19/07/2023. In the further course of the transit until 20/07/2023, the wind came from different directions with 2-3 Bft, also due to local pressure systems. On the morning of 20/07/2023, RV METEOR reached the main research area near the coast of western Sicily as planned.

High pressure continued to dominate the weather. The initially weak to moderate westerly to north-westerly wind increased to 5-6 Bft in the afternoon, the wind sea was 0.5 m. Until 23/07/2023, an undisturbed high with weak pressure contrasts prevailed over the Mediterranean. With a significant wave height of less than 0.5 m and a mostly light to moderate southerly wind, sometimes veering, research could be carried out without restrictions. In the morning of 21/07/2023, fog patches appeared in the humid air mass near the coast of Sicily.

From 24/07/2023 onwards, a weak depression formed over northern Italy and moved eastwards. The high south of Sicily briefly intensified, temporarily increasing the wind force to 5 Bft. During the night of 25/07/2023, the sirocco briefly reached the working area, leaving dusty residues on the ship and instruments. Night temperatures reached 33°C. On 25/07/2023, the weather was again calm. The south-easterly wind of 4 to 5 Bft decreased temporarily during the day. In the afternoon, the wind veered northwesterly and increased steadily to 6-7 Bft during the night. In the morning of 26/07/2023, a weak cold front from the depression over northern Italy crossed the area of research. The significant wave height increased post-frontal until the evening to 2 m. The weak cold front moved into a high over the Gulf of Gabes and the Ionian Sea and weakened further.

From 27/07/2023, a ridge of high pressure over the Azores extended into the western Mediterranean and affected the working area on its eastern side. The significant wave height decreased to 0.5 m. Southerly to southeasterly winds of 4 Bft persisted until 29/07/2023. The wind then decreased, causing fog to develop in the working area during the night from 30/07/2023 to the morning. During the day, the wind veered northwest and increased to 3-4 Bft. However, the significant wave height was only 0.5 m. At the vicinity of the Azores high, which stretched from Morocco to Tunisia, the wind increased to 5-6 Bft from northwesterly directions on 31/07/2023, with a significant wave height of 1.5 m. The wind speed decreased to 3 Bft during the night until the morning of 01/08/2023. In the afternoon, the wind increased again, reaching 5 to 6 Bft from

northwest, with a significant wave height of 1.5 m. This situation continued until the end of the day on 01/08/2023. This weather situation and its effects continued until the end of the research work on 02/08/2023.

The weather for the transit to Piraeus was characterized by a weak high-pressure influence and a low-pressure system over northern Italy with low-pressure differences over the Ionian Sea. RV METEOR reached the port of Piraeus in the morning of 05/08/2023, with mostly light westerly and later southerly winds from the Aegean, and a significant wave height around 0.5m.

# **7 Station List M191**







# **8 Data and Sample Storage and Availability**

(J. Geldmacher, A. Micallef, G.M. Ferrante, O.A. Archontikis)

The rock samples recovered during cruise M191 will be stored at the rock repository at GEOMAR Helmholtz Centre for Ocean Research Kiel. They will be analysed at the home institutions of the participating scientists and possibly further cooperating institutions, and the obtained analytical results will be published in English language in peer-reviewed journals and thus made publicly available. Upon request, individual samples will be made available to third parties after analysis, data interpretation and publication. Filter membranes of seawater samples will be stored at the facilities of NHM, London and resulting datasets from analyses will be made available after peerreviewed publication. The bathymetric and sediment echo sounding data as well as the sound probe data will be archived in the IT storage infrastructure at GEOMAR. Magnetic field data will be stored at National Institute of Oceanography and Applied Geophysics (OGS), Trieste (with a copy at GEOMAR).

Data sharing and exchange will take place within the Ocean Science Information System (OSIS) maintained by the GEOMAR data management team. Bathymetric raw data are submitted to the Federal Maritime and Hydrographic Agency (Bundesamt für Seeschifffahrt und Hydrographie, BSH). Processed magnetic data will be stored at the World Data Center PANGAEA, bathymetric

data will be uploaded to the World Data Center PANGAEA and the International Hydrographic Organization Data Centre for Digital Bathymetry (IHO DCDB). For a three-year moratorium, however, the high-resolution bathymetric data from the working area will be available to the project members only. A 100 m grid will be submitted to EMODnet Bathymetry at the end of 2024. Availability of the geochemical and micropaleontological data is restricted until publication.

<b>Type</b>	<b>Database</b>	Available	<b>Free Access</b>	Contact
M <sub>191</sub> metadata	OSIS	Oct. 2023	Oct. 2023	jgeldmacher@geomar.de
Rock sample data	OSIS, Georoc, PetDB	n/a	After publication	jgeldmacher@geomar.de
Echo-sounding data (working area) (KONGSBERG EM 122, EM710, PARASOUND)	BSH, OSIS, <b>PANGAEA</b>	Oct. 2023	Oct. 2026	jgeldmacher@geomar.de aaron.micallef@um.edu. mt
Sound probe data (XSV- 02)	BSH, OSIS, <b>PANGAEA</b>	Oct. 2023	Nov. 2023	jgeldmacher@geomar.de
Magnetic field data (Magnetometer)	<b>PANGAEA</b>	Oct. 2023	After publication	elodolo@ogs.it, gferrante@ogs.it
Seawater sample data	<b>PANGAEA</b>	After publication	After publication	odysseas.archontikis@un iv.ox.ac.uk

**Table 8.1 Overview of data availability**

# **9 Acknowledgements**

First of all, we would like to thank Master Derk Apetz and his skilful crew for their excellent support, professionalism and for providing a very pleasant working atmosphere on board, which contributed enormously to the success of this expedition. We also thank Andreas Raeke and Holger Jens from the German National Meteorological Service (DWD) for the daily weather briefings. We are grateful to Danilo Cavallaro for providing existing multi-beam bathymetric data of Terribile Bank, and Dario Civile and Emanuele Lodolo (both OGS) for being part of the SUAVE project team and their ongoing shore-based support. Filippo Muccini (INGV, La Spezia) is especially thanked for providing the magnetometer which was of vital importance for localizing and identifying the possible presence of volcanic outcrops. Maik Lange, Mia Schumacher, Bettina Domeyer, Anke Bleyer and Karin Junge (all GEOMAR) are thanked for help with cruise preparations at GEOMAR. Special thanks are due to Jens O. Herrle (Goethe University Frankfurt) for kindly providing the equipment and supplies for filtration of seawater, and his encouragement. We are also grateful to the Geschäftsstelle des Gutachterpanels Forschungsschiffe (GPF), the Leitstelle Deutsche Forschungsschiffe (LDF) and Briese Research for their support. We are very grateful to Briese Research and the LDF for obtaining detailed data on the distribution of submarine cables in the working area, without which dredging at most positions could not have been carried out. Last but not least, we would like to thank the countries of Spain, Italy, Malta and Greece for permitting our research. This cruise was funded by the German Research Foundation (Deutsche Forschungsgemeinschaft, DFG) and the GEOMAR Helmholtz Centre for Ocean Research Kiel. Aaron Micallef was supported by the David and Lucile Packard Foundation.

# **10 References**

- Archontikis, O. A., & Young, J. R., 2021. A reappraisal of the taxonomy and biodiversity of the extant coccolithophore genus Palusphaera (Rhabdosphaeraceae, Prymnesiophyceae). Phycologia, 60(6), 589–602.
- Archontikis, O. A., Millán, J. G., Andruleit, H., Cros, L. Kleijne, A., Heldal, M., Doan-Nhu, H., Winter, A., Blanco-Bercial, L. & Young, J. R., 2023a. Taxonomy and morphology of Calciopappus curvus sp. nov. (Syracosphaeraceae, Prymnesiophyceae), a novel appendagebearing coccolithophore, Protist.
- Archontikis, O. A., Millán, J. G., Winter, A., & Young, J. R., 2023b. Taxonomic re-evaluation of Ericiolus and Mercedesia (Prymnesiophyceae) and description of three new species. Phycologia, 62(2), 179–193.
- Archontikis, O. A., Young, J. R., & Cros, L., 2020. Taxonomic revision and classification of extant holococcolithophores previously placed in the genus Anthosphaera Kamptner emend. Kleijne 1991. Acta Protozoologica, 59(3–4).
- Armitage, J. J., Ferguson, D. J., Goes, S., Hammond, J. O. S., Calais, E., Rychert, C. A., & Harmon, N., 2015. Upper mantle temperature and the onset of extension and break-up in Afar, Africa. Earth and Planetary Science Letters, 418, 78-90.
- Beccaluva, L., Colantoni, P., Di Girolamo, P., & Savelli, C., 1981. Upper-Miocene submarine volcanism in the Strait of Sicily (Banco senza Nome). Bulletin of Volcanology, 44.
- Calanchi, N., Colantoni, P., Rossi, P. L., Saitta, M., & Serri, G., 1989. The Strait of Sicily continental rift systems: Physiography and petrochemistry of the submarine volcanic centres. Marine Geology, 87, 55-83.
- Calo, M., & Parisi, L., 2014. Evidences of a lithospheric fault zone in the Sicily Channel continental rift (southern Italy) from instrumental seismicity data. Geophysical Journal International, 199, 219-225.
- Caracausi, A., Favara, R., Italiano, F., Nuccio, P. M., Paonita, A., & Rizzo, A., 2005. Active geodynamics of the central Mediterranean Sea: Tensional tectonic evidences in western Sicily from mantle-derived helium. Geophysical Research Letters, 32(4).
- Cavallaro, D., and Coltelli, M., 2019. The Graham Volcanic Field offshore Southwestern Sicily revealed by high-resolution seafloor mapping and ROV images. Frontiers in Earth Science 7, 311.
- Civile, D., Lodolo, E., Accaino, F., Geletti, R., Schiattarella, M., Giustiniani, M., et al., 2018. Capo Granitola-Sciacca Fault Zone (Sicilian Channel, Central Mediterranean): Structure vs magmatism. Marine and Petroleum Geology, 96, 627-644.
- Civile, D., Lodolo, E., Zecchin, M, Ben-Avraham, Z., Baradello, L., Accettella, D., Cova, A., Caffau, M., 2015. The lost Advernture Adchipelago (Sicilian Channel, Mediterranean Sea): Morpho-bathymetry and Late Quaternary palaeogeographic evolution. Global and Planetary Change 125, 36-47.
- Civile, D., Lodolo, E., Alp, H., Ben-Avraham, Z., Cova, A., Baradello, L., et al., 2014. Seismic stratigraphy and structural setting of the Adventure Plateau (Sicily Channel). Marine Geophysical Research, 35(1), 37-53.
- Civile, D., Lodolo, E., Tortorici, L., Lanzafame, G., & Brancolini, G.,2008). Relationships between magmatism and tectonics in a continental rift: The Pantelleria Island region (Sicily Channel, Italy). Marine Geology, 251, 32-46.
- Coltelli, M., Cavallaro, D., D'Anna, G., D'Alessandro, A., Grassao, F., Mangano, G., Patane, D., Gresta, S., 2016. Exploring the submarine Graham Bank in the Sicily Channel. Annals of Geophysics 59(2), S0208.
- Cros L. & Fortuño J.-M., 2002. Atlas of northwestern Mediterranean coccolithophores. Scientia Marina 66 (Suppl. 1): 7–182.
- D'Amario, B., Ziveri, P., Grelaud, M., Oviedo, A. & Kralj, M., 2017. Coccolithophore haploid and diploid distribution patterns in the Mediterranean Sea: can a haplo-diploid life cycle be advantageous under climate change? Journal of Plankton Research. **9**: 1–14.
- Falzone, G., Lanzafame, G., Rossi, P.L., 2009. Il vulcano Ferdinandea nel Canale di Sicilia. Geoitalia 29, 15-20.
- Fedorik, J., Toscani, G., Lodolo, E., Civile, D., Bonini, L., & Seno, S., 2018. Structural analysis and Miocene-to-Present tectonic evolution of a lithospheric-scale, transcurrent lineament: The Sciacca Fault (Sicilian Channel, Central Mediterranean Sea). Tectonophysics, 722, 342- 355.
- Ferguson, D. J., Barnie, T. D., Pyle, D. M., Oppenheimer, C., Yirgu, G., Lewi, E., et al., 2010. Recent rift-related volcanism in Afar, Ethiopia. Earth and Planetary Science Letters, 292(3- 4), 409-418.
- Francis, P., 1995. Fire and water. Geology Today Jan-Feb 1995, 27-31.
- Franke, D., 2013. Rifting, lithosphere breakup and volcanism: comparison of magma-poor and volcanic rifted margins. Marine and Petroleum Geology, 43, 63-87.
- Ghisetti, F. C., Gorman, A. R., Grasso, M., & Vezzani, L., 2009. Imprint of foreland structure on the deformation of a thrust sheet: The Plio-Pleistocene Gela Nappe (southern Sicily, Italy). Tectonics, 28, TC4015.
- Kendall, J. M., Stuart, G. W., Ebinger, C. J., Bastow, I. D., & Keir, D., 2005. Magma assisted rifting in Ethiopia. Nature, 433, 146-148.
- Kleijne, A., 1993. Morphology, Taxonomy and distribution of extant coccolithophorids (Calcareous nannoplankton). Drukkerij FEBO B.V., Katwijk. 1–321.
- Knappertsbusch, M., 1993. Geographic distribution of living and Holocene coccolithophores in the Mediterranean Sea. Marine Micropaleontology. **21**: 219–247.
- Lodolo, E., Civile, D., Zanolla, C., Geletti, R., 2012. Magnetic signature of the Sicily Channel volcanism. Marine Geophysical Research 33, 33-44.
- Lodolo, E., Civile, D., Zecchin, M., Sante Zampa, L., Accaino, F., 2019. A series of volcanic edifices discovered a few kilometers off the coast of SW Sicily. Marine Geology 416, 105999.
- Oviedo, A., Ziveri, P., Álvarez, M. & Tanhua, T., 2015. Is coccolithophore distribution in the Mediterranean Sea related to seawater carbonate chemistry? Ocean Science. **11**: 13–32.
- Putirka, K., & Platt, B., 2012. Basin and Range volcanism as a passive response to extensional tectonics. Geosphere, 8(6), 1274-1285.
- Romagnoli, C., Belvisi, V., Innangi, S., Di Martino, G., Toniello, R., 2020. New insights on the evolution of the Linosa volcano (Sicily Channel) from the study of its submarine portions. Marine Geology 419, 106060.

Rotolo, S. G., Castorina, F., Cellura, D., & Pompilio, M., 2006. Petrology and geochemistry of submarine volcanism in the Sicily Channel Rift. The Journal of Geology, 1414, 355-365.

Siebert, L., Simkin, T., 2023. Global Volcanism Program (www.volcano.si.edu)

Washington, H.S., 1909). The submarine eruptions of 1831 and 1891 near Pantelleria. American Journal of Science 27 (158), 131.

### **11 Appendices**

- **11.1 M191 Dredge Station Details and Rock Description (including graphical rock descriptions)**
- **11.2 Underway Water Sampling List for Nannoplankton Research**

TS: thin section billet **Amph:** amphibole MI: melt inclusions MI: melt inclusions CHEM: chemistry slab to prepare materials for geochemical analysis Apt: apatite Apt: manganese Mn: manganese Ar/Ar: estimate of sample quality for <sup>40</sup>Ar<sup>39</sup>Ar dating Bi: biotite Bi: biotite Bi: biotite Mt: magnetite Mt: magnetite Gl/MIN: potential glass and / or mineral separates Cc: calcite Cc: calcite CC: calcite Cl: olivine SED: sediment Chl: chlorite Opx: orthopyroxene REF: reference sample for immediate transport to home institution Cpx: clinopyroxene Pl: plagioclase Pl: plagioclase after cruise Factors of the cruise of th

#### **Abbreviations in Table Header: Abbreviations for Minerals and Materials:**

Gc: geochemistry Gz: quartz<br>
Gather Company Company<br>
Company C  $Gm:$  ground mass Ilm: ilmenite Cc: carbonate

Pp: phosphate





#### **M191-8-1**

#### **SSE slope of cone from base to top (Area 1 seamount 3)**



Comments: several very large coral/carbonate encrusted blocks. Open textured irregular to platy forms. 3 clasts have smoother rounder forms, their sawn surface shows encrusted clastic rock with angular volcanic fragments

















**Area 2, seamount 2. Repeat of 16 on opposite side. Northern flank of seamount as described at dredge site 16**





#### **M191-20-1**

#### **Area 2, Largest volcano on Graham Bank. Sampling of NE flank of volcano, dredging up flank that appears to have had a mass wasting event**

















#### **M191-24-1**

#### **Area 2 seamount 4, cone base diameter 675m, 190-80m SW flank from base to top**









Comments: carbonate sample with many displaying small volcanic clasts within the softer calcareous muddy parts



#### **Area 2, northernmost volcano on Graham Bank**











#### **M191-42-1**

8. Secondary Minerals: n/a

on outer surface 10. Comment:

9. Encrustations: encrusted in biogenic material

### **Western working area onTerribile bank, prominent oval-shaped cone rising over small basin; SW facing slope**





**M191-44-1**

<u>Dredge on bottom UTC, hrs, °N, °E, depth may be a manuscus of the state of the state of the state of the state o</u><br>Dredge off bottom UTC, hrs, °N, °E, depth may be a state of the state of the state of the state of the sta Dredge off bottom UTC, hrs, °N, °E, depth m 19:02 1/8 full **Western working area ofTerribile bank, NW-SE elongated high with area of high reflection in backscatter. Dredge across high** total volume:

Comments: biogenic blocks, rhodoliths and encrusted material - all carbonate









#### **M191-49-1**










#### **M191-55-1**



10. Comment: Description and interpretation on

attached sketch















10. Comment:



















#### **Nameless Bank, NE facing, middle part** Dredge on bottom UTC, hrs, °N, °E, depth m 12:19 36°51,89'N 13°06,54'E 495 Dredge off bottom UTC, hrs, °N, °E, depth m 13:35 36°51,85'N 13°06,52'E 430 total volume: Few rocks Two small dense carbonate fragmentsNOTES Gl/MIN CHEM Ar/Ar SAMPLE # SAMPLE DESCRIPTION TS PICTURE 이ᇛ M191-75-1-1 1. Rock Type: Hard very dense limestone, white to Similar location to dredge 49, plus similar recovery Similar location to dredge 49, plus similar recovery grey to pale grey with infill of micritic mud, vein like texture 2. Size: 10.5x7x4.5cm 3. Shape / Angularity: soubrounded to subangular 4. Color of cut surface: Brown to organe exterior, white to light grey interior with dark brownish veining  $M191 - 75 - 1 - 1$ 5. Texture / Vesicularity: Massive with some veining  $\mathcal{S}$  , where  $\mathcal{S}$  $\overline{a}$ 6. Phenocrysts: n/a 7. Matrix: Cement, well indurated hardground 8. Secondary Minerals: n/a 9. Encrustations: minor encrustations 10. Comment: Branching hard coral also found alongside the rocks, with encrusted specimens, see sample 2.

**M191-75-1**



















**M191-84-1**













Dense lava block, broken form, plus two large open textured coral









## Sample #: M191-13-1-1

[PLEASE FILL OUT]

### TO BE APPENDED TO SAMPLE DESCRIPTION SHEET



### Sample #: M191-20-1-1

[PLEASE FILL OUT]

### TO BE APPENDED TO SAMPLE DESCRIPTION SHEET



### Sample #: M191-20-1-2 Result TO 4. [PLEASE FILL OUT]

#### TO BE APPENDED TO SAMPLE DESCRIPTION SHEET

 $\frac{1}{2}$ 



FRUM HOLE

### Sample #: M191-34 -1- 14 / 18 [PLEASE FILL OUT]

 $\mathbf{r}$ 

#### TO BE APPENDED TO SAMPLE DESCRIPTION SHEET



NOTE: ANE BLACA ENCASIVE COVID BE GREAMIC MATERIA.



 $\left( \frac{\partial}{\partial t} \right)$ 

 $(2)$ 

PYROCURSTIC DENSITY CURRENT:

- "PLUCKING UP BLUCKS / PEBBLES OF COUNTRY ROCK MATERIAL (NEAR STORE ENVIRONMENT)
- · PUMICES / OCUPTION EITHER INTERSECTED A MIDROTHERAM FIELD OF INCORPORATION WARR FROM THE SURROWANTS ENVICONMENT (SEAMATER?). This COWERED EMPIREMENT , POLDERATURES + INCREASED THE WATER CONTENT LEADING TO ALADAMIN OF THE PULLES 10 ZEOLINE (?)
- . FLOW MM HAVE BEEN SUBMACINE, BUT IT IS LIKELY THIS REPRESENTE A SUSABRIAL OFPOSIT. COLUMN

Deposit PEGON? PEBBLES OF ERODED COUNTRY ROCK IN LORPORTED  $S<sub>FA</sub>$ 

Fru DEPOSIT FROM EVOLVED EXUPTION FROM PONTELERIA, LINGSA, ETNA?:

. It is possible THAT THIS IS NOT FROM NAMELESS BANK ITSELF, BUT REPRESENS THE FRIENT FROM A MITCH ECUPRON FROM A MITCH (WORK) VOLONIC CEMPRE.

Ly EXECAMINAUS WOULD BE PHAT!

A) IT was BE FINCR GUAINED (BUT IT MAY BC ON THE DISTORANTL  $A9.3$ 

 $MWO$ 

(B) WOULD BE FUND ELECTRICE - E.G. AT GRAMMAN BANK. . THIS WOULD HAVE BEEN ZEOLITISED STARTING AFTER POPOSITION.

NOTE: THE PRESENCE OF PUMILE (MICROVENTR) MORE EVOLVED, L'HEMICALLY) cirsts is workenive of a money Sion conduct ERUPTION. MOLE EVALUED MARCHA IS CONGERLY MORE EXPLOSIVE, POSSIBLE THAT IT IS UNISUR MEE OF ERVARAN...? E.G. PUMICE CONE VOLCANISM?  $AN$ 

## Sample #: M191- 37-1-1

### [PLEASE FILL OUT]

TO BE APPENDED TO SAMPLE DESCRIPTION SHEET

FREVERSE HAS POSSIBLE INTERPRETATIONS



# **Sample #: M191-49-1-3** [PLEASE FILL OUT]

 $\overline{9}$ 

 $\sqrt{2}$ 




# Sample #: M191- 55-1-

# [PLEASE FILL OUT]

#### TO BE APPENDED TO SAMPLE DESCRIPTION SHEET



INTERPETATION:

NOTE: THIS SUTTONS,  $PHTO5,$ LECTRE + SMALES FOR AUMUDIS. PEPERATE CONTACT/Cammonstap BENJERS MACHA (LANA) + WET SEDMONT. SOLVENT IS CALCAREDUS IN CURPOSITION, PERIMPS LAJA OVERCOPE THE SEDIMENT OR SQUEEZED OUT INTO IT. THIS MM EXPLAIN THE PEPERINE CONTROL ON ONE SIDE OF THE BUCK? RIM IS SURROUNDING WHOLE SERVINGS, SUCRESNAG TOSTACE IS Perman.

# **Sample #: M191- 83-1-**  $\frac{1}{5}$

[PLEASE FILL OUT]

#### TO BE APPENDED TO SAMPLE DESCRIPTION SHEET



- (1) PEPERINE TEADLE FROM INTRUSIVE/EMPLACED MEMATIC BODY (E.G. LAJA) WITH ORANGE SEPTIMENT BETO BAROO SEDS? NO CLASSY KIMS ANE TO SEPTIMENT CONTACT.
- (2) MYMOCASTITE FORM IN SITU BEECUTION OF THE MARMATIC MATERIAL & TINGE 'SCRIMENT' REINE ASTOCHED GUISS + PALMONIC FOSING THE MEMORE APPOINTER IN PLACE. Q: WHERE ARE THE GASS PLAIS IN THIS MODEL ?

### **Appendix 11.2**

# **Underway Water Sampling List for Nannoplankton Research**

<b>Sample</b>	<b>Date</b>	Time (UTC)	<b>Depth</b> (m)	Latitude (°N)	Longitude (°E)	Tempera- ture $(^{\circ}C)$	<b>Salinity</b> (PSU)
M191-PL-1	16.07.23	08:45:20	$0-5$	36°03,086'	$-05^{\circ}22,111'$	19.2	
M191-PL2	16.07.23	12:35:48	$0-5$	36°05,054'	$-04^{\circ}38,177'$	24.6	36.35
M191-PL2b	16.07.23	12:52:00	$0-5$	36°05,345'	$-04^{\circ}35.683'$	24.8	36.35
M191-PL3	16.07.23	18:20:11	$0 - 2$	36°12,678'	$-03°31,374'$	20.8	36.33
M191-PL4	16.07.23	22:16:20	$0 - 2$	36°18,041'	$-02^{\circ}46,118'$	26.4	36.74
M191-PL5	16.07.23	02:05:15	$0 - 2$	36°24,450'	$-02^{\circ}01,598'$	27.5	36.83
M191-PL6	16.07.23	06:21:55	$0 - 2$	36°39,327'	$-01^{\circ}12,518'$	27.1	36.96
M191-PL7	16.07.23	10:23:55	$0 - 2$	36°59,058'	00°02,977'	27.3	37.12
M191-PL8	16.07.23	14:16:30	$0 - 2$	37°21,397'	00°08,706'	27.8	36.97
M191-PL9	16.07.23	17:54:17	$0 - 2$	37°41,589'	00°45,490'	28	36.74
M191-PL10	16.07.23	21:42:33	$0 - 2$	$37^\circ 46'$	01°24'	28.3	37.32

**Water samples taken within the Spanish EEZ:**

# **Water samples taken within the Italian EEZ:**





		Time	Depth	Latitude	Longitude	Tempera-	Salinity
Sample	Date	(UTC	(m)	(°N)	(°E)	ture (°C)	(PSU)
M191-PL63	3.8.23	15:11	$0 - 2$	36°12,883'	018°28,390'	29.8	37.68
M191-PL64	3.8.23	6:21	$0 - 2$	36°11,511'	019°06,533'	28.9	38.47
M191-PL65	3.8.23	20:59	$0 - 2$	36°10,403'	019°37,530'	28.7	38.81
M191-PL66	3.8.23	23:09	$0 - 2$	36°09,498'	020°02,901'	28.6	39.05
M191-PL67	4.8.23	0:53	$0 - 2$	36°08,769'	020°23,200'	28.4	38.99
M191-PL68	4.8.23	3:03	$0 - 2$	36°07,918'	020°46,928'	28.2	38.42
M191-PL69	4.8.23	5:00	$0 - 2$	36°07,112'	021°09,405'	27.8	38.64
M191-PL70	4.8.23	7:07	$0 - 2$	36°06,209'	021°35,098'	28.2	39.03
M191-PL71	4.8.23	9:07	$0 - 2$	36°05,318'	021°59,349'	27.7	38.63
M191-PL72	4.8.23	11:08	$0 - 2$	36°04,352'	022°23,161'	27.4	38.94
M191-PL73	4.8.23	13:12	$0 - 2$	36°03,579'	022°47,888'	28.2	39.35
M191-PL74	4.8.23	14:48	$0 - 2$	36°03,755'	023°06,881'	27.6	38.96
M191-PL75	4.8.23	17:14	$0 - 2$	36°22,521'	023°20,736'	27.9	39.02
M191-PL76	4.8.23	19:01	$0 - 2$	36°38,638'	023°26,700'	28.4	38.94
M191-PL77	4.8.23	21:10	$0 - 2$	36°57,852'	023°33,223'	28	39.33
M191-PL78	4.8.23	23:03	$0 - 2$	37°14,587'	023°38,176'	28.3	38.7
M191-PL79	5.8.23	1:02	$0 - 2$	37°31,848'	023°43,174'	28.2	38.88
M191-PL80	5.8.23	3:02	$0 - 2$	37°45,333'	023°42,544'	28.4	38.97

**Water samples taken within the Greek EEZ:**